

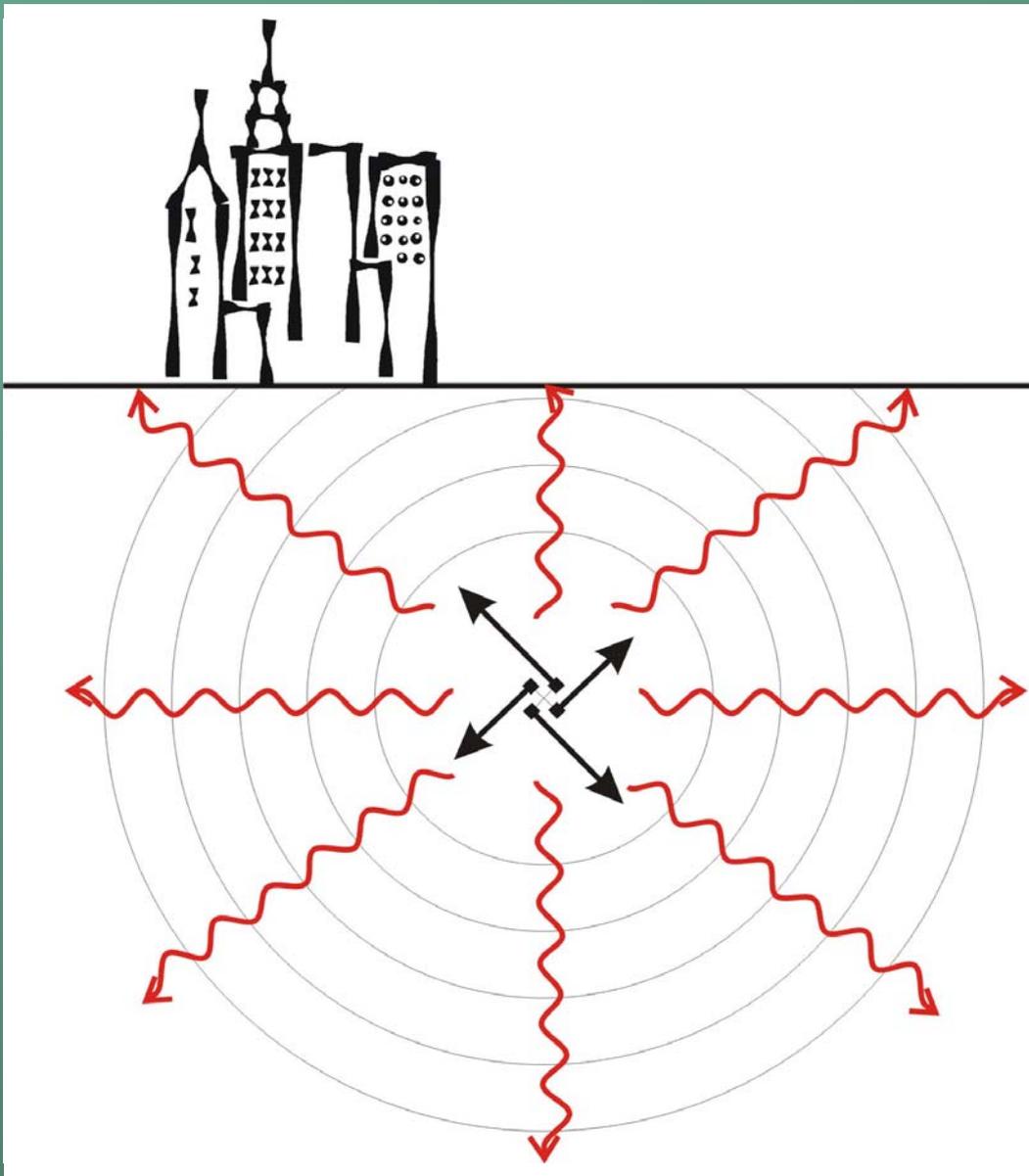
Theoretical Seismology 2: Wave Propagation

Based on a lecture by James Mori of the
Earthquake Hazards Division,
Disaster Prevention Research Institute,
Kyoto University



Contents

- Rays
Snell's Law
Structure of the Earth
- Seismic Waves
Near-Field Terms (Static Displacements)
Far-Field Terms (P, S, Surface waves)
- Normal modes
Free oscillations of the Earth



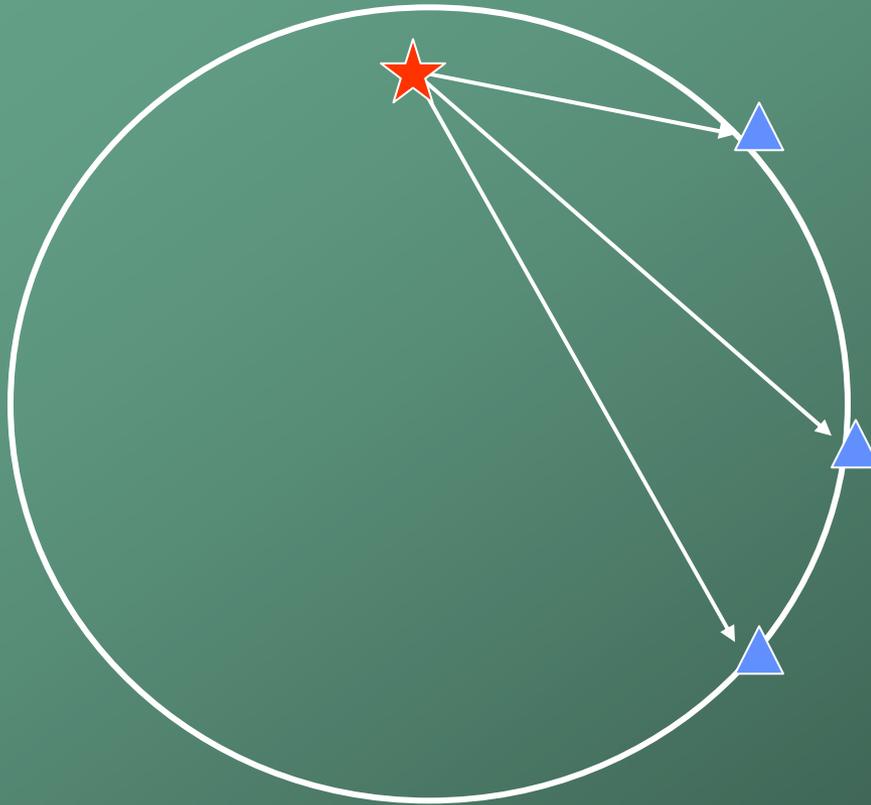
Faulting



Seismic waves



Homogeneous Earth

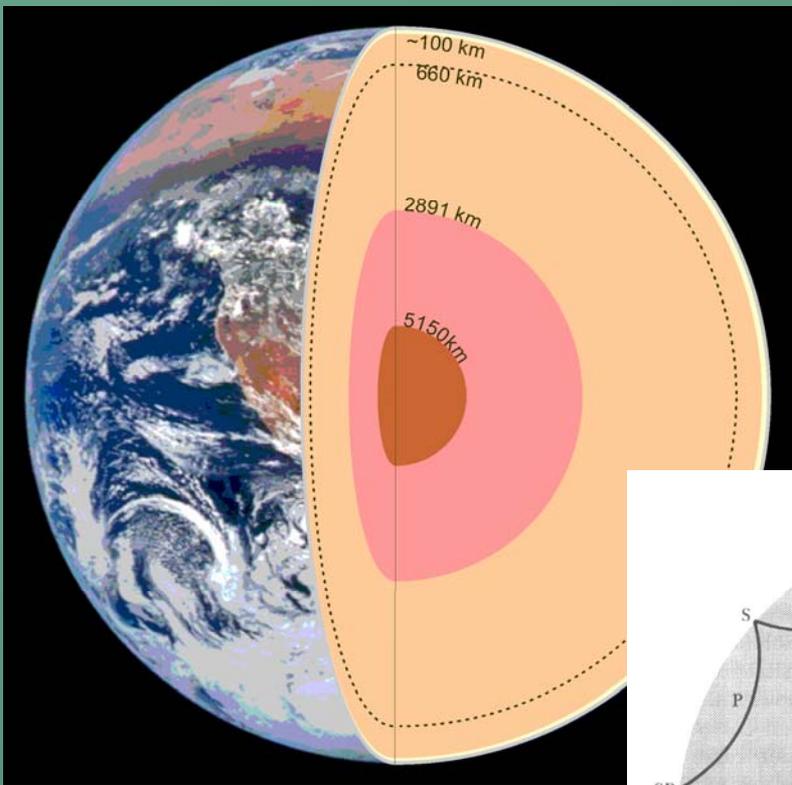


If the Earth had constant velocity the wave paths would be very simple.

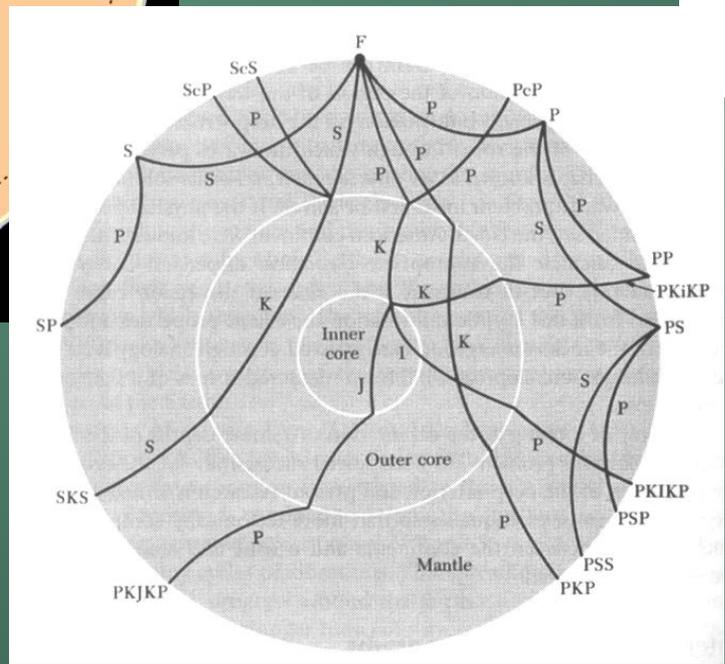
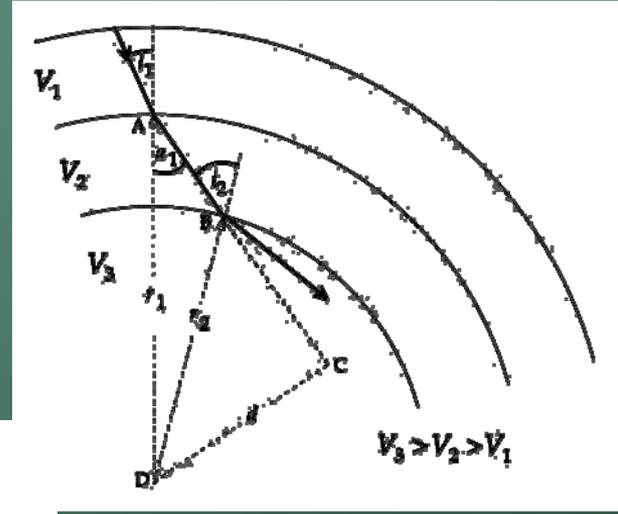


Structure in the Earth results in complicated paths

Lowrie, 1997, fig 3.69



USGS

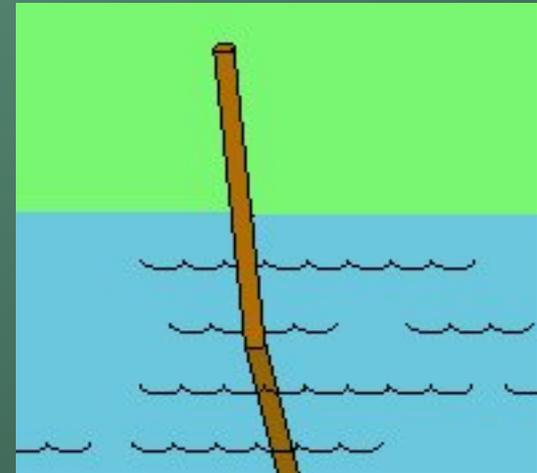
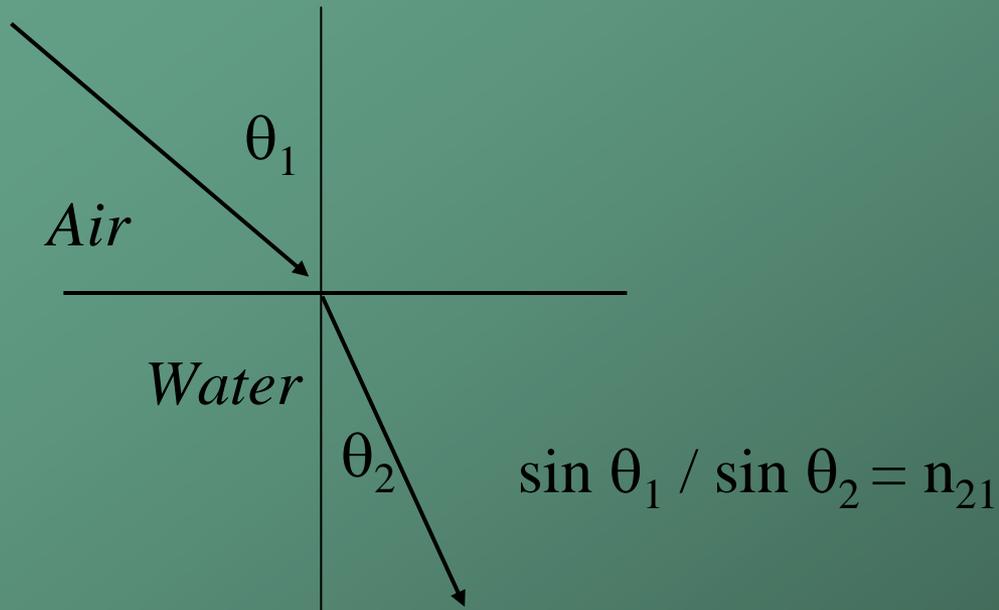


Bolt, 2004, fig 6.3



Rays

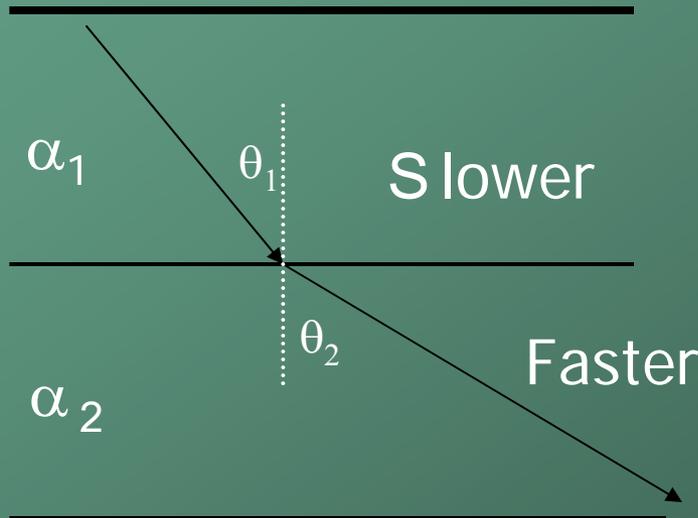
Snell's Law Fermat's Principle



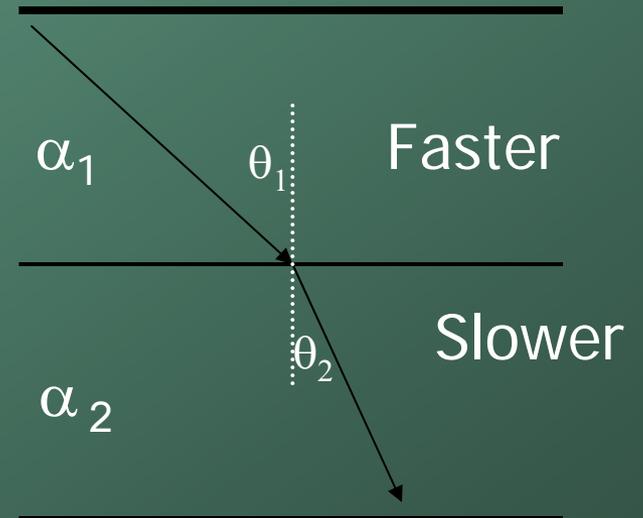


Ray Paths in a Layered Medium

α = velocity of seismic energy in the layer



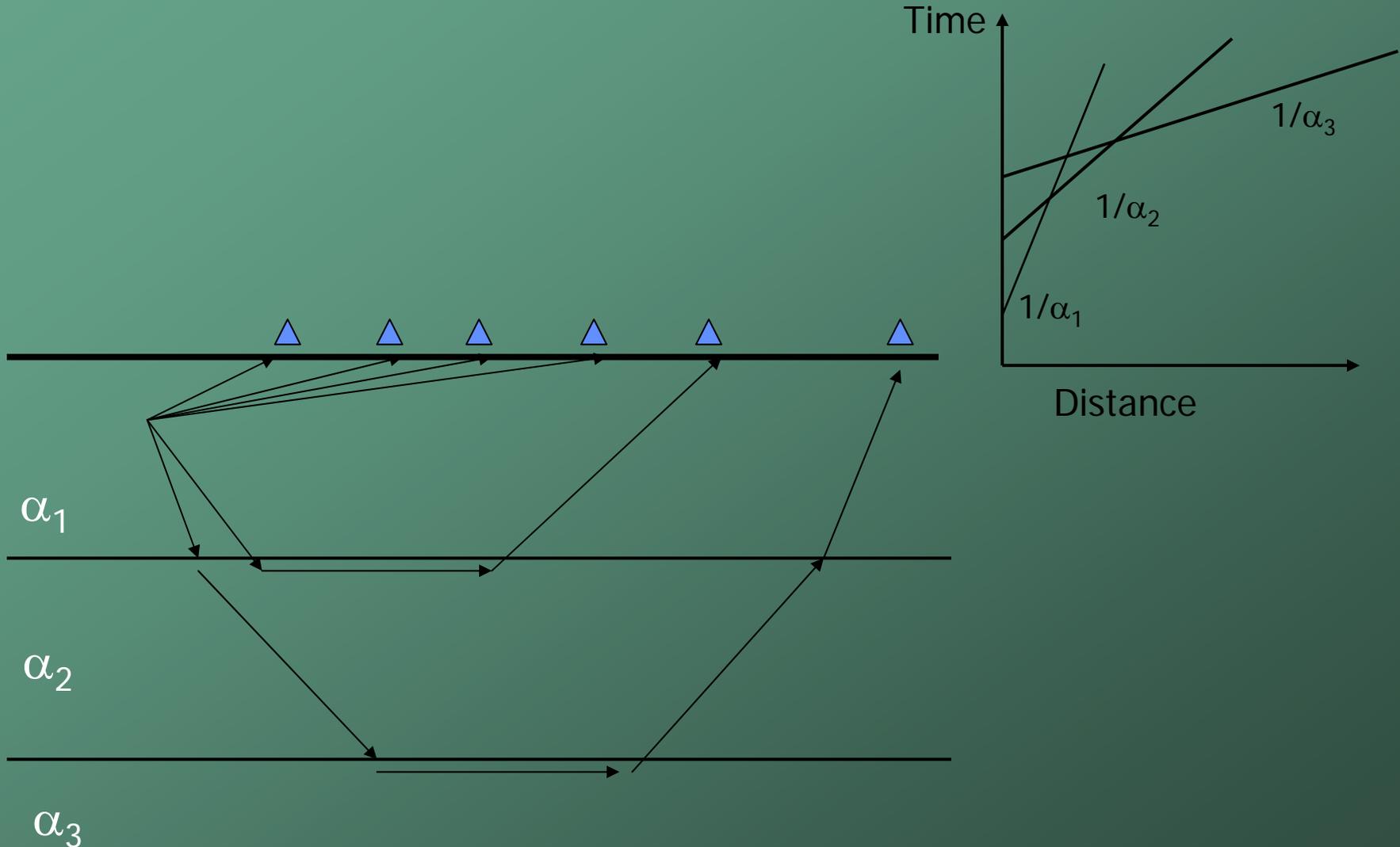
$$\alpha_1 < \alpha_2$$



$$\alpha_1 > \alpha_2$$



Ray Paths in a Layered Medium

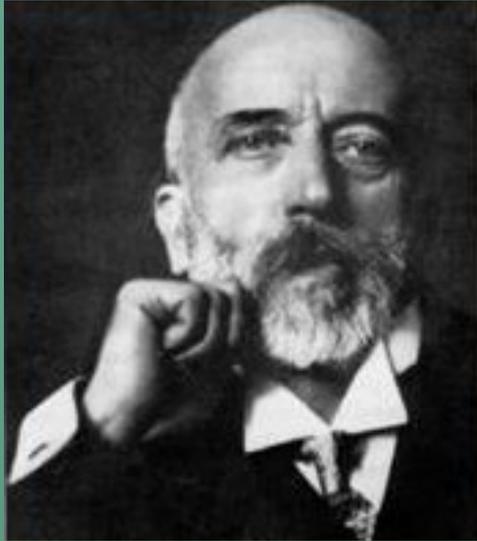


α_3



The Moho

Andrija Mohorovicic (1857-1936)

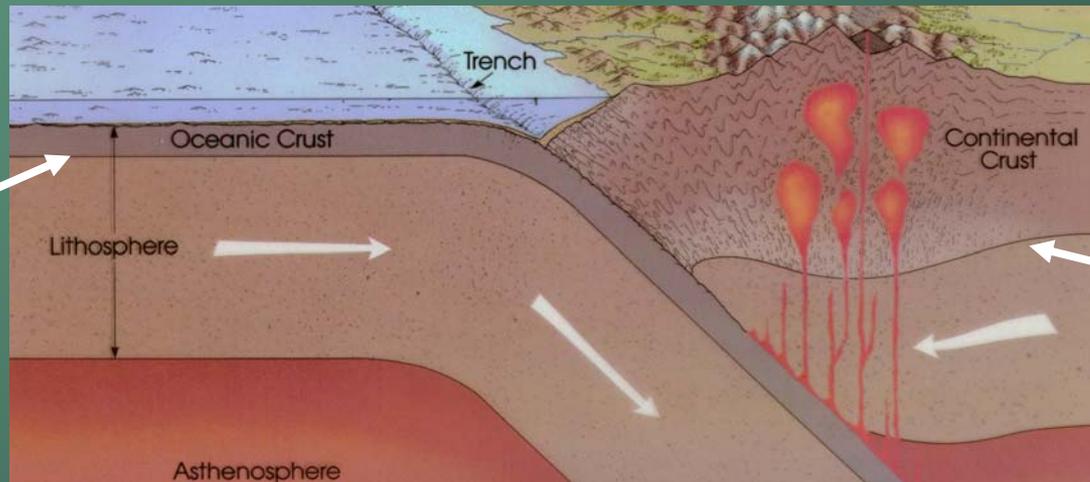


Found seismic discontinuity at 30 km depth in the Kupa Valley (Croatia).

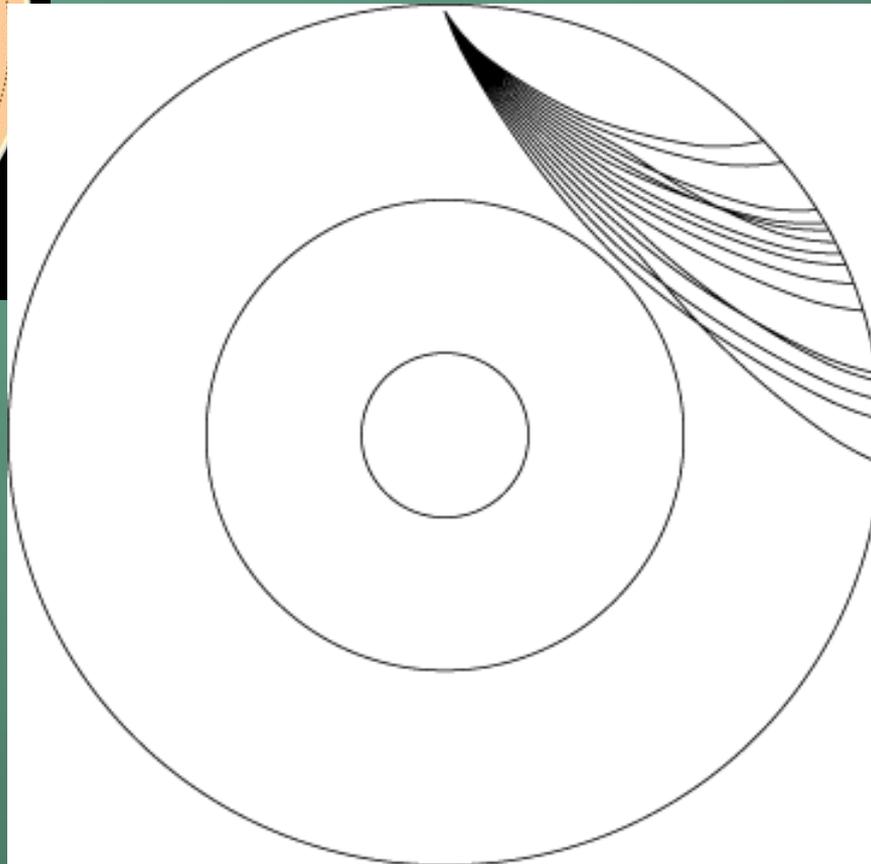
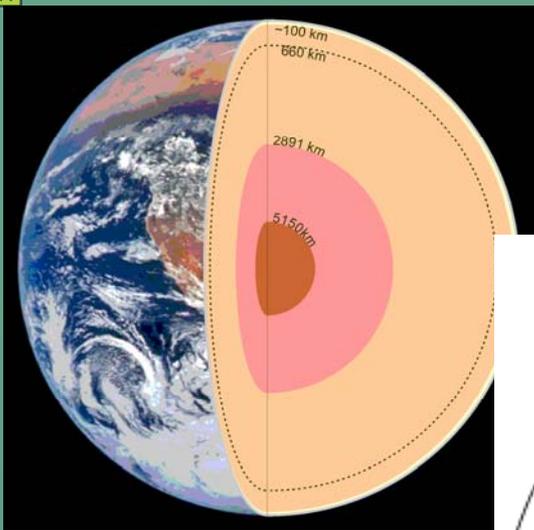
Mohorovicic discontinuity or 'Moho'

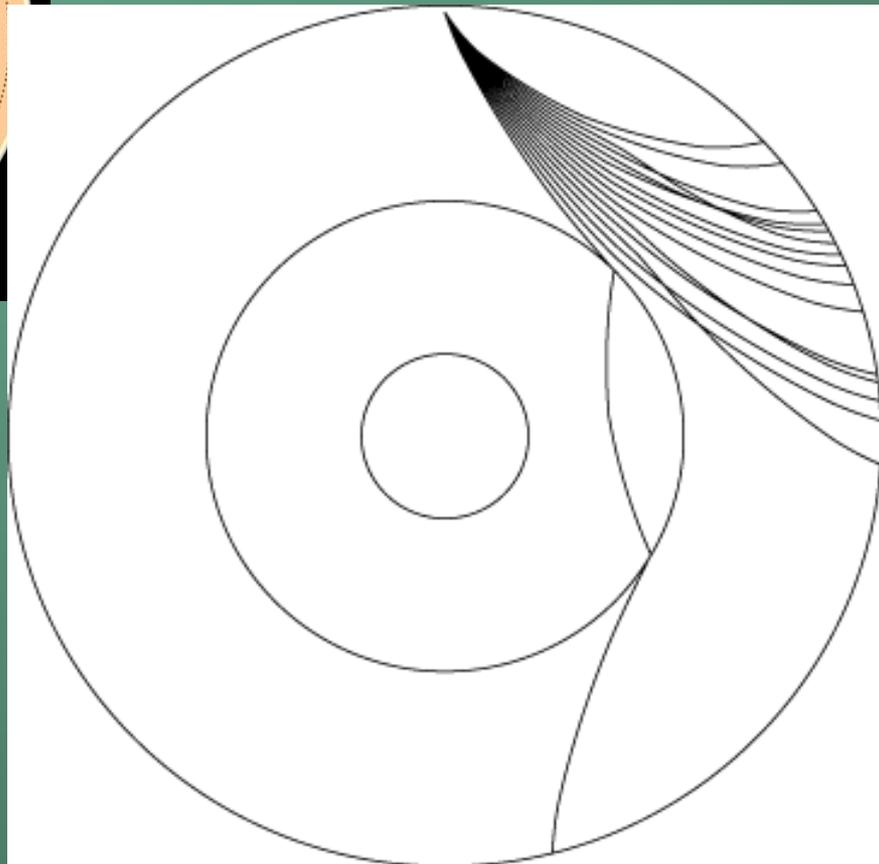
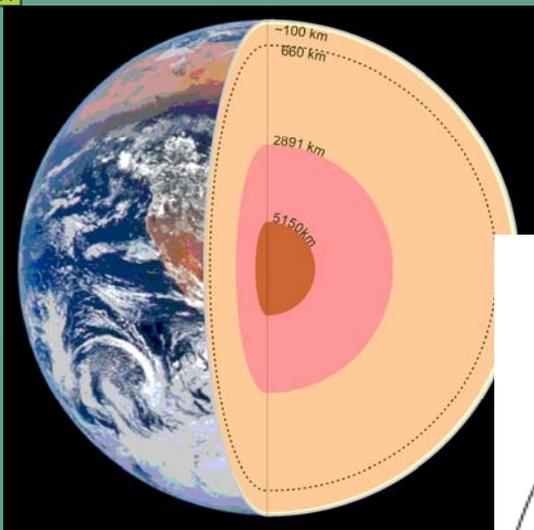
Boundary between crust and mantle

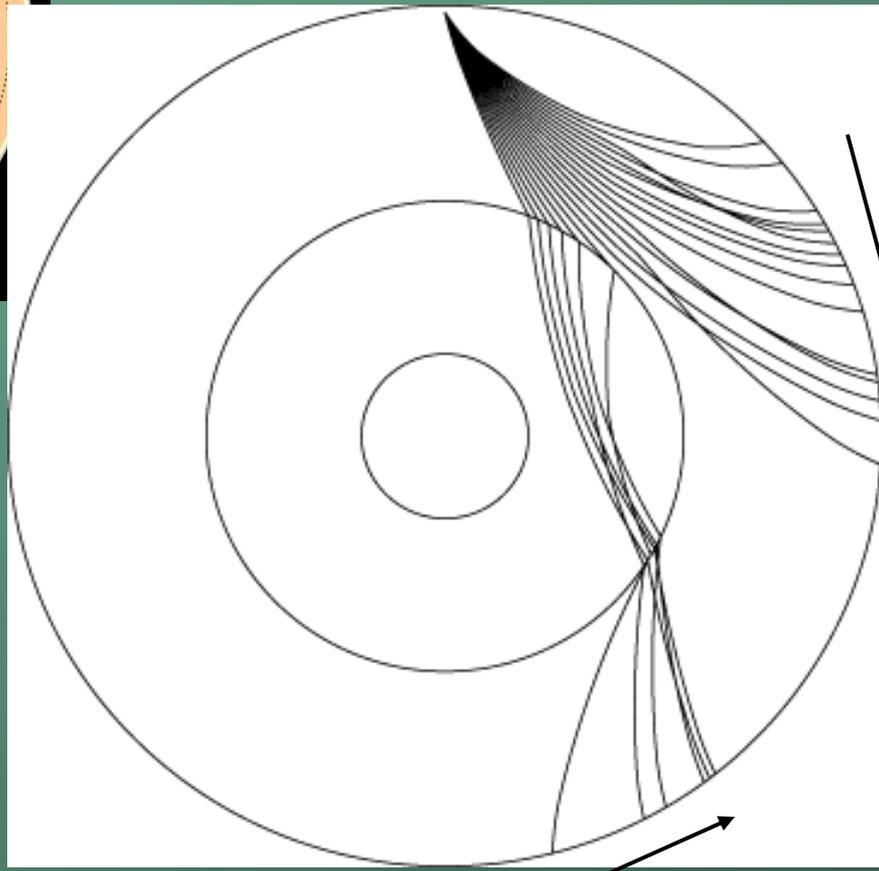
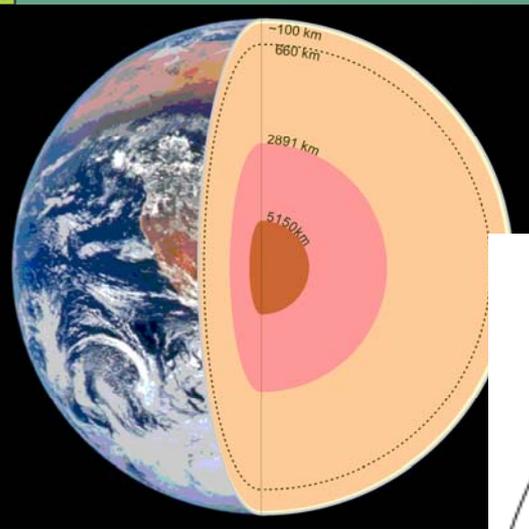
The Moho



The Moho

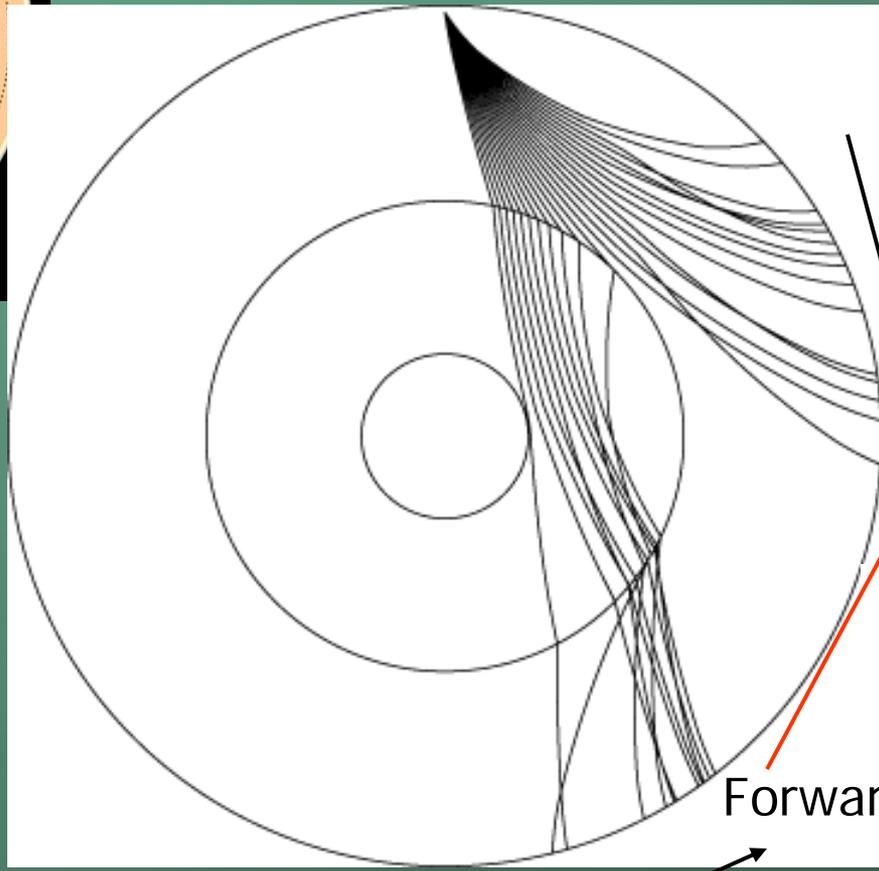
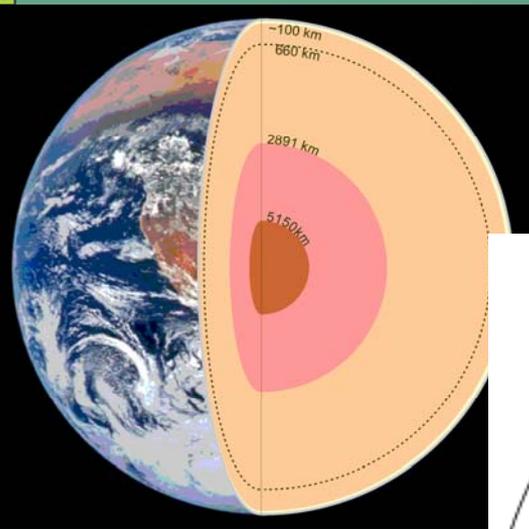






Forward Branch

Backward Branch

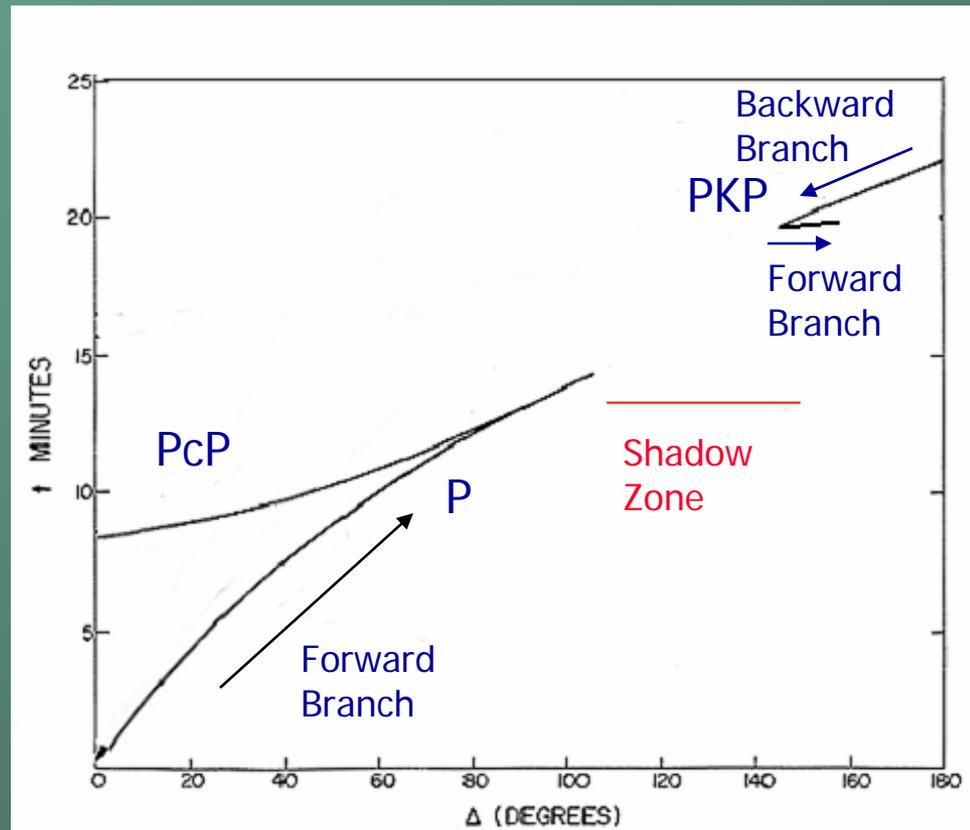
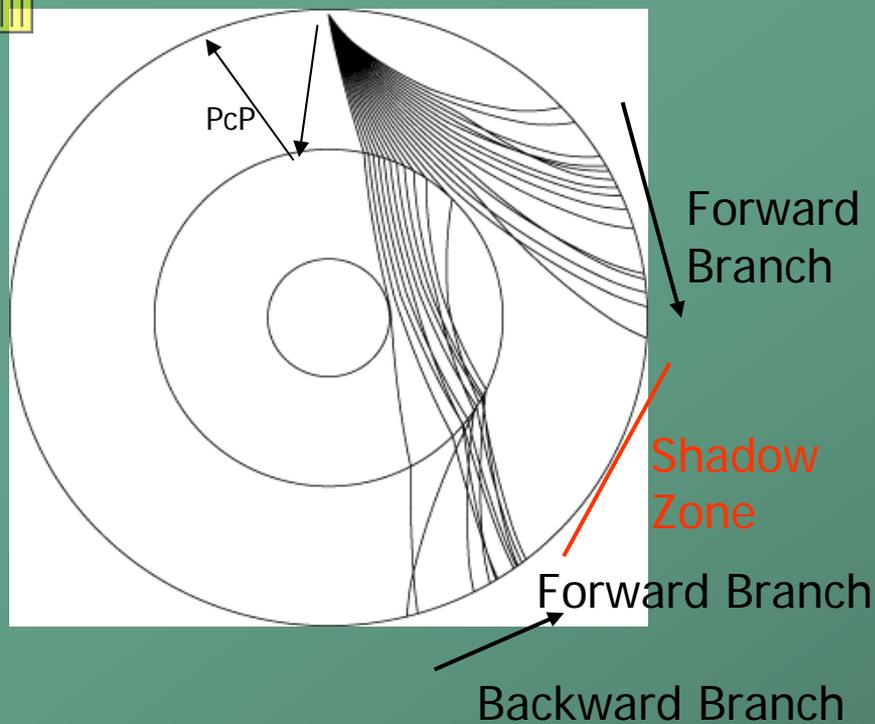


Forward Branch

Shadow Zone

Forward Branch

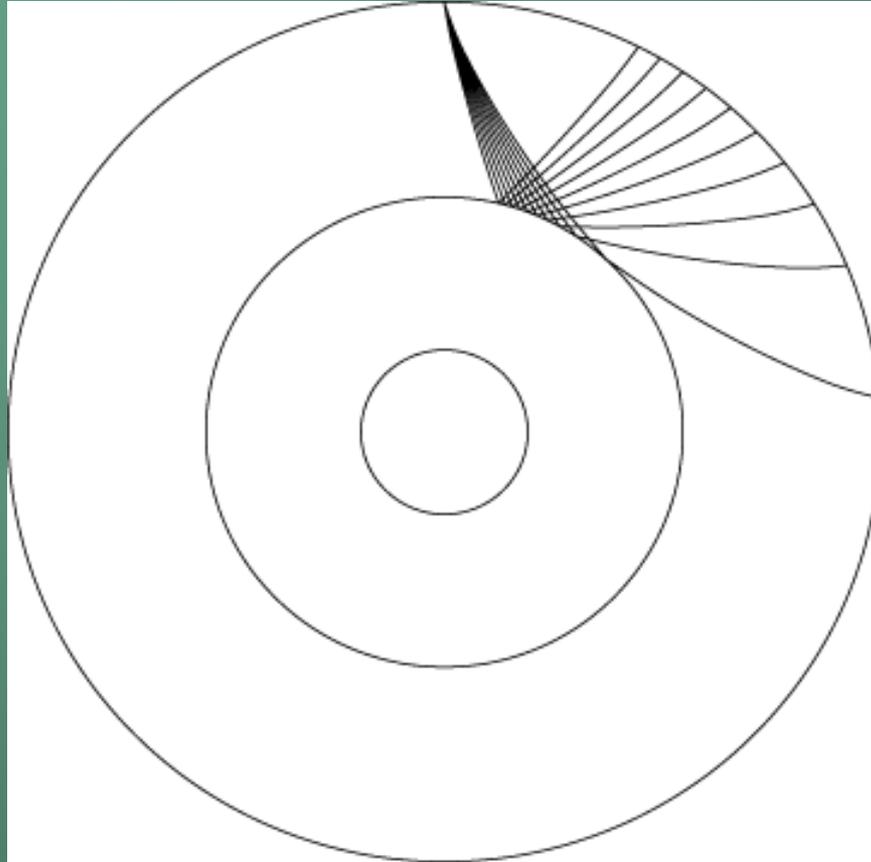
Backward Branch



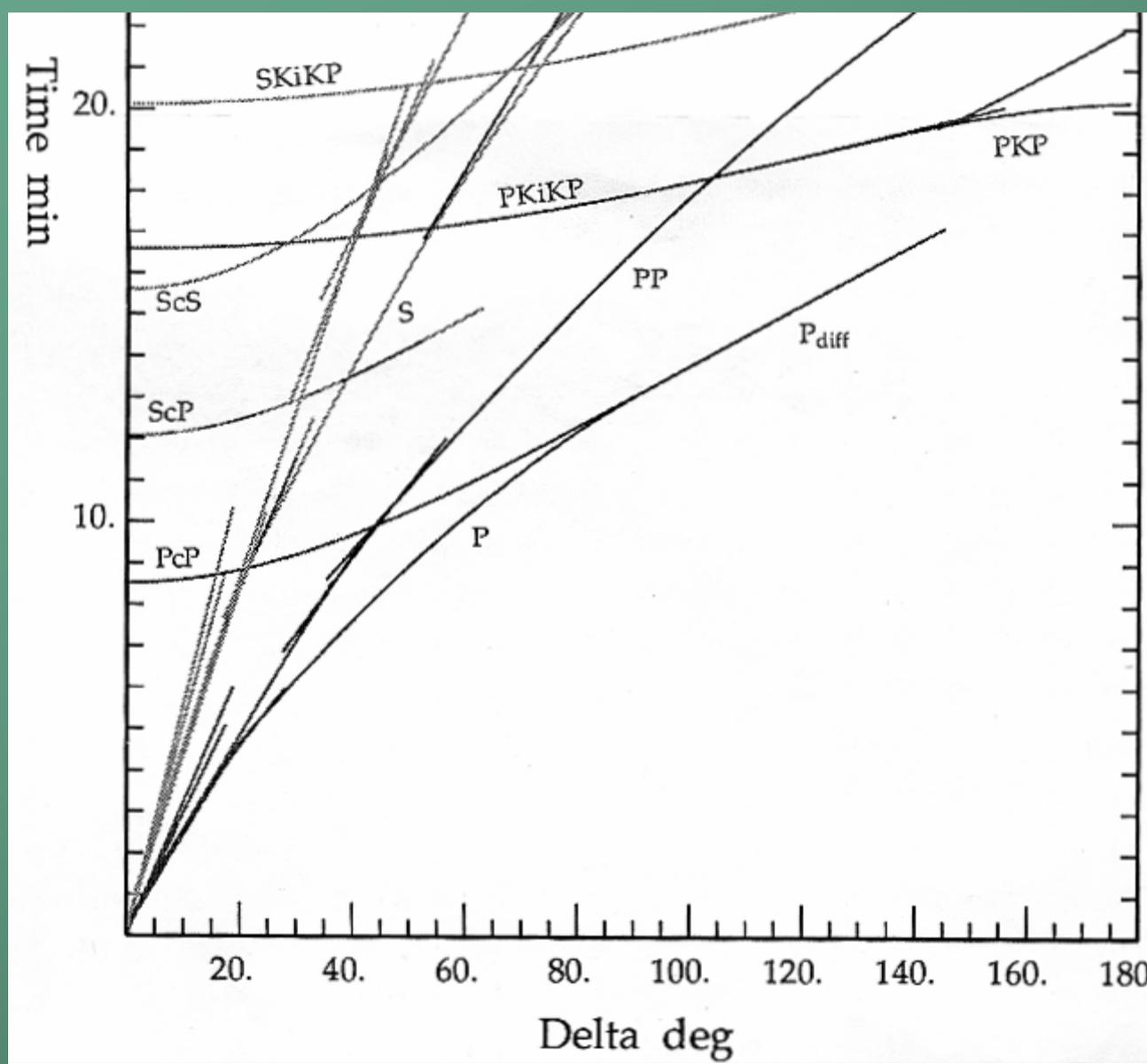
- 1912 Gutenberg observed shadow zone 105° to 143°
- 1939 Jeffreys fixed depth of core at 2898 km (using PcP)



PcP



Core Reflections



P	Mantle P
S	Mantle S
K	Outer core P
I	Inner core P
c	Reflection from the outer core
i	Reflection from the inner core
diff	Diffracted arrival

IASP91, Kennett and Engdahl, 1991



Seismic Waves

Aspects of Waves not Explained by Ray Theory

- Different types of waves (P, S)
- Surface Waves
- Static Displacements
- Frequency content

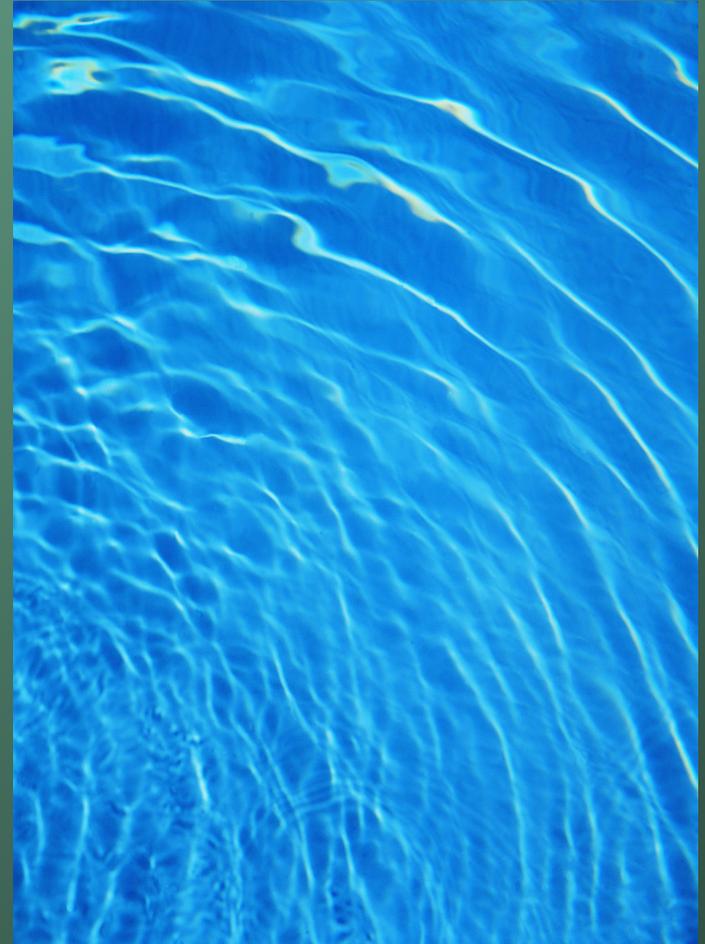
Wave Equation

$$\frac{\partial^2 u_1}{\partial x_1^2} = \frac{1}{c} \frac{\partial^2 u_1}{\partial t^2}$$

1-D wave equation

c = propagation speed

This is the equation that explains the waves on a spring: constant velocity wave propagation, no mass transfer, different from circulation eq.





1-D Wave Equation

$$\frac{\partial^2 u_1}{\partial x_1^2} = \frac{1}{c} \frac{\partial^2 u_1}{\partial t^2}$$

Solution

$$u(x, t) = A \sin[\omega(t \pm x / c)]$$

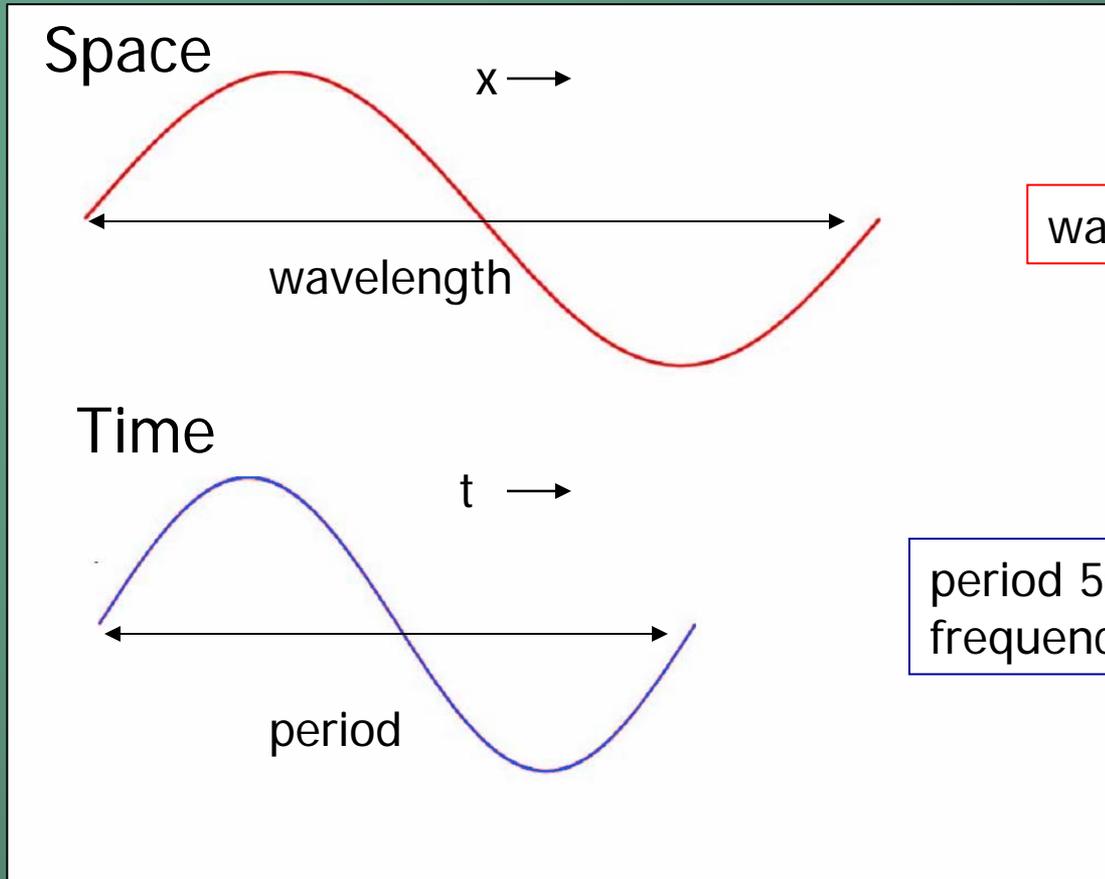
$$T = \frac{2\pi}{\omega}$$

T = wave period

ω = angular frequency

Wave Period and Wavelength

Velocity 6 km/s



wavelength 300 km

period 50 s
frequency = $1/\text{period} = 0.02$ hz

$$\text{Velocity} = \text{Wavelength} / \text{Period}$$

Period

Wavelength

Body waves (P· S)	0.01 to 50 sec	50 m to 500 km
Surface waves	10 to 350 sec	30 to 1000 km
Free Oscillations	350 to 3600 sec	1000 to 10000 km
Static Displacements	∞	-

3-D Wave Equation with Source

$$\rho \frac{\partial^2 u}{\partial t^2} = f + (\lambda + 2\mu) \nabla(\nabla \cdot u) - \mu \nabla \times (\nabla \times u)$$

source

spatial 2nd derivative

Solution

$$u(x, t) = \frac{1}{4\pi\rho} A^N \frac{1}{r^4} \int_{r/\alpha}^{r/\beta} \tau \cdot M_0(t - \tau) d\tau + \frac{1}{4\pi\rho\alpha^2} A^{IP} \frac{1}{r^2} M_0\left(t - \frac{r}{\alpha}\right) + \frac{1}{4\pi\rho\beta^2} A^{IS} \frac{1}{r^2} M_0\left(t - \frac{r}{\beta}\right) + \frac{1}{4\pi\rho\alpha^3} A^{FP} \frac{1}{r} \dot{M}_0\left(t - \frac{r}{\alpha}\right) + \frac{1}{4\pi\rho\beta^3} A^{FS} \frac{1}{r} \dot{M}_0\left(t - \frac{r}{\beta}\right)$$

Near-field Terms (Static Displacements)

Far-field Terms (P, S Waves)

Near-field terms

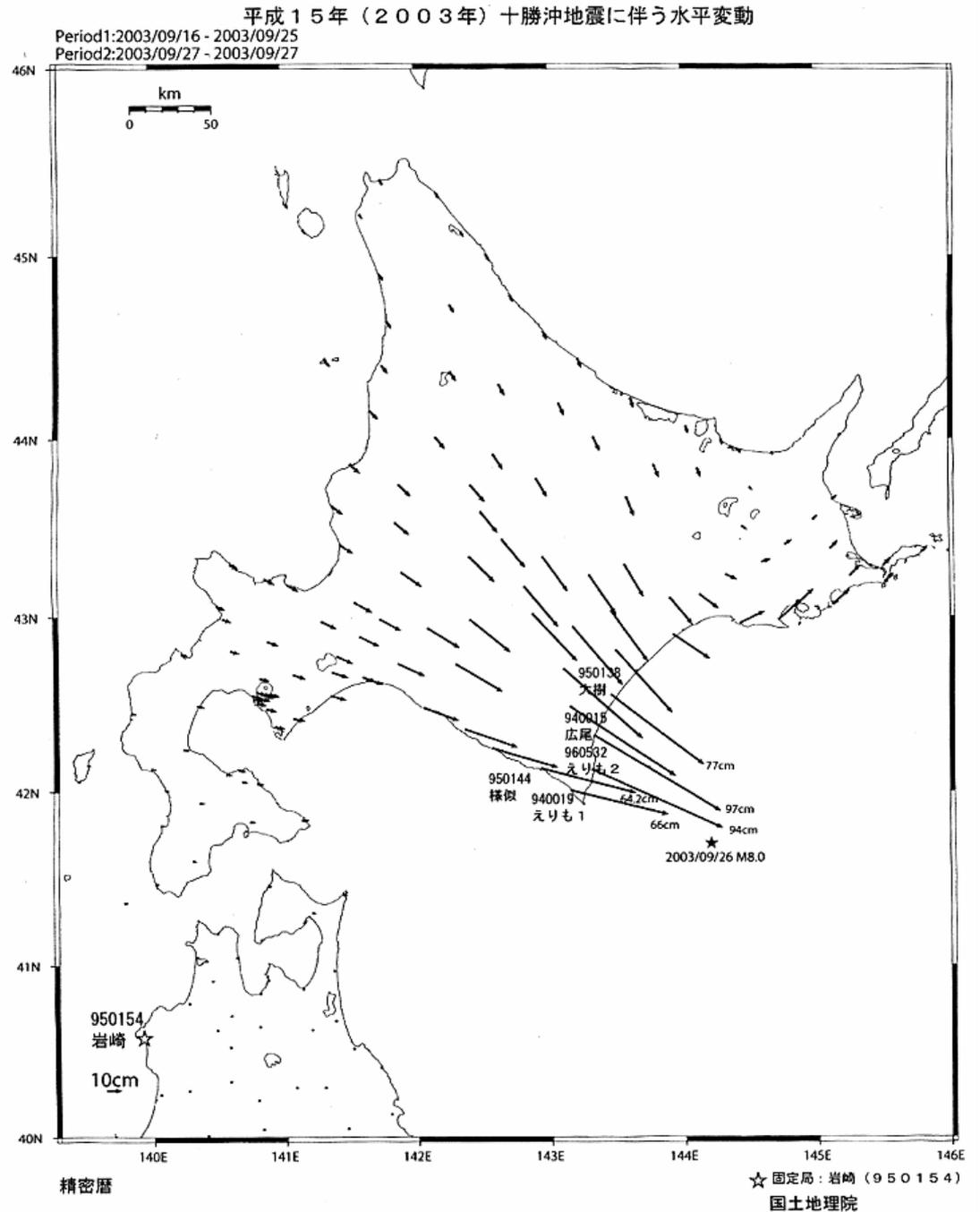
- Static displacements
- Only significant close to the fault
- Source of tsunamis



Bei-Fung Bridge near Fung-Yan city, 1999 Chi-Chi, Taiwan earthquake

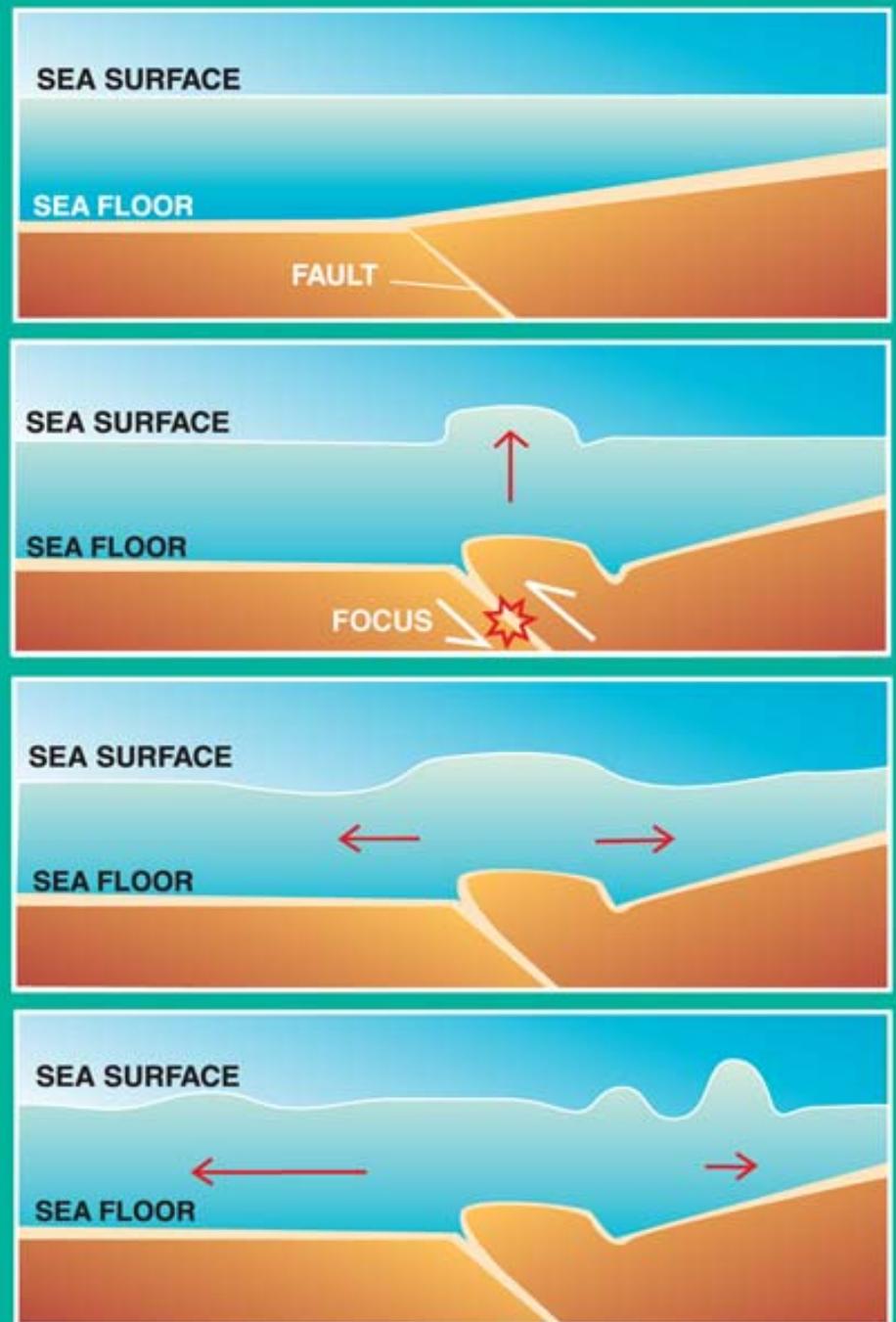
Static displacements

Co-seismic deformation of 2003 Tokachi-oki Earthquake (M8.0)



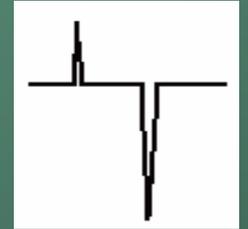


Generation of Tsunami from Near-field Term

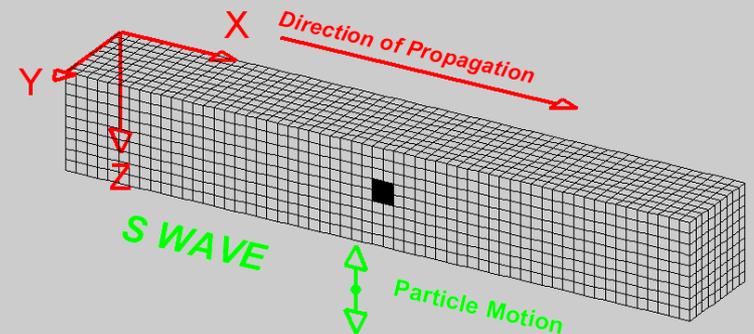
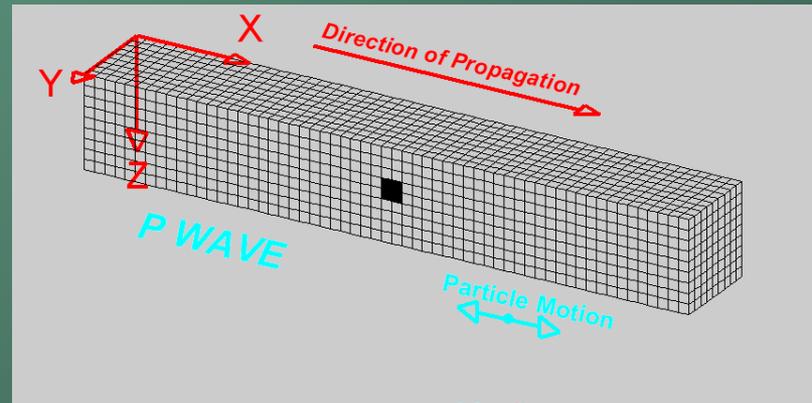


Far-field Terms

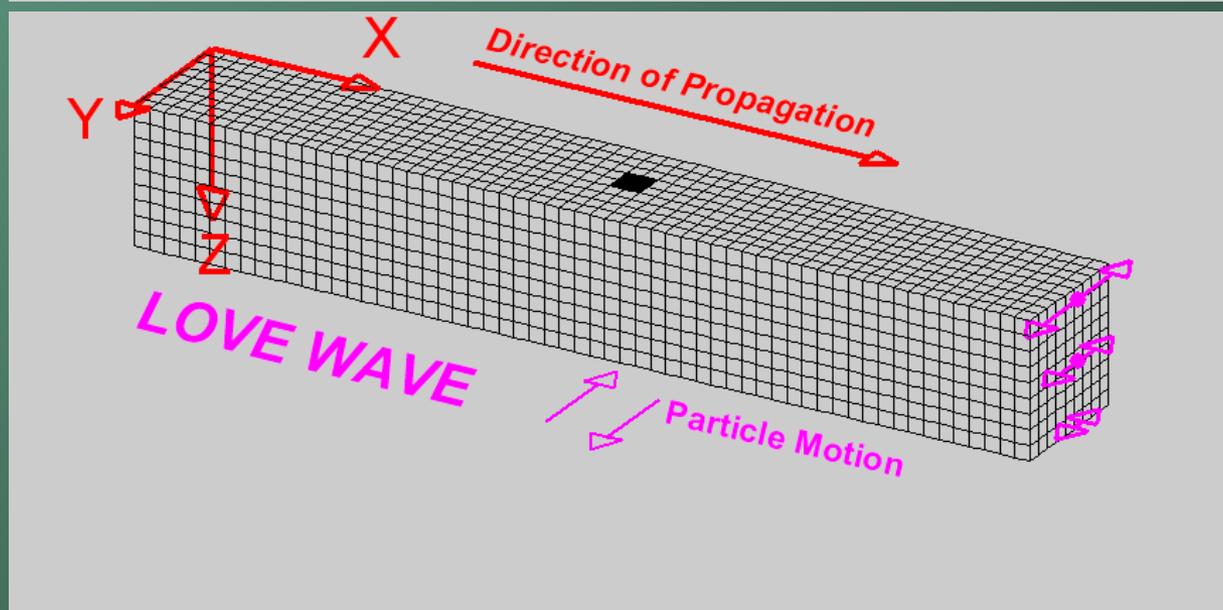
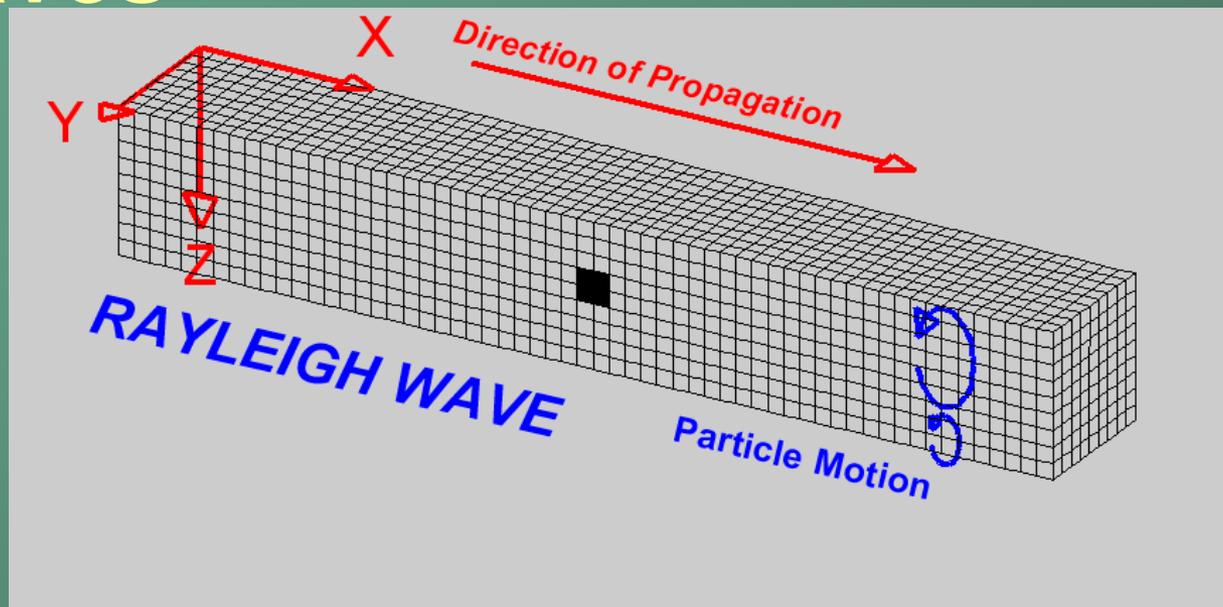
$$+\frac{1}{4\pi\rho\alpha^3}A^{FP}\frac{1}{r}\dot{M}_0\left(t-\frac{r}{\alpha}\right)+\frac{1}{4\pi\rho\beta^3}A^{FS}\frac{1}{r}\dot{M}_0\left(t-\frac{r}{\beta}\right)$$



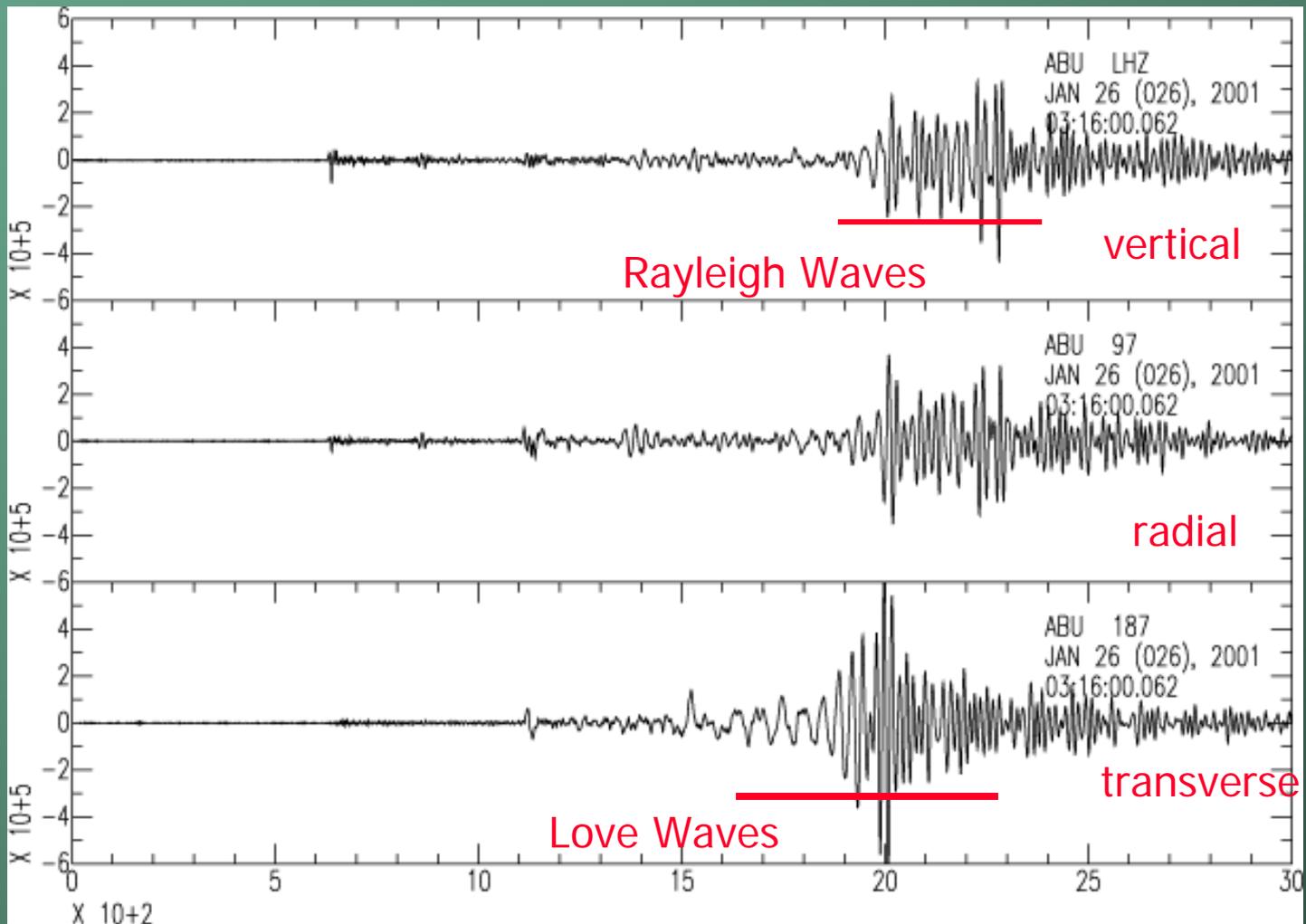
- Propagating Waves
- No net displacement
- P waves
- S waves



surface waves



January 26, 2001 Gujarat, India Earthquake (Mw7.7)



Recorded in Japan at a distance of 57° (6300 km)

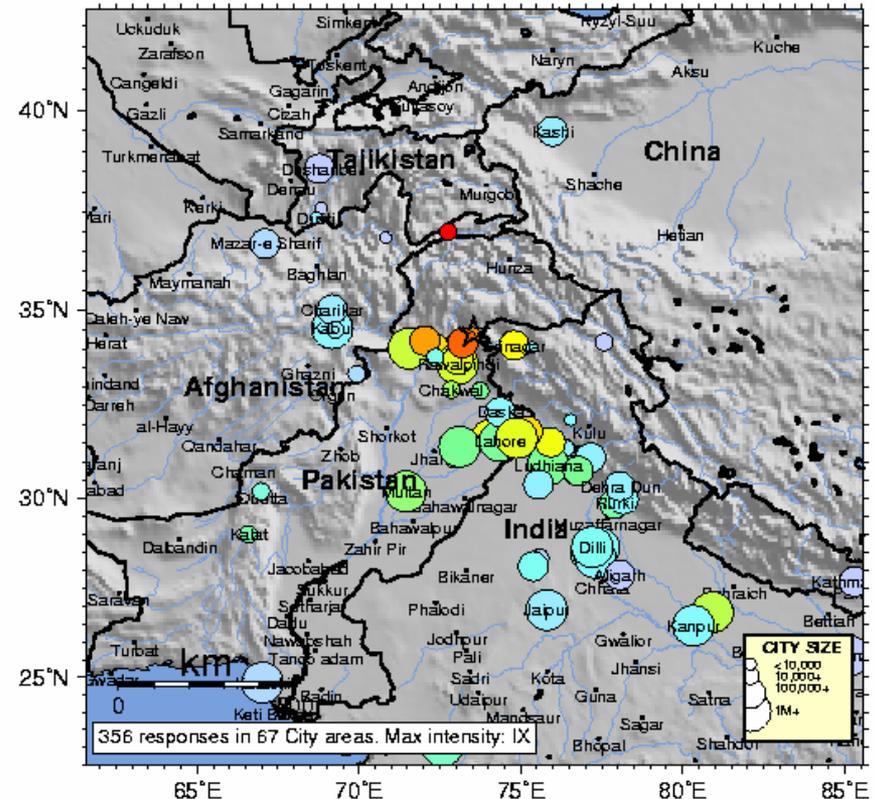
Amplitude and Intensity

Seismic waves lose amplitude with distance traveled - attenuation

$$A(t) = A_0 e^{-\omega_0 t / 2Q}$$

So the amplitude of the waves depends on distance from the earthquake. Therefore unlike magnitude intensity is not a single number.

USGS Community Internet Intensity Map (58 miles NNE of Rawalpindi, Pakistan)
ID:dyae_05 03:50:39 GMT OCT 08 2005 Mag=7.6 Latitude=N34.43 Longitude=E73.54



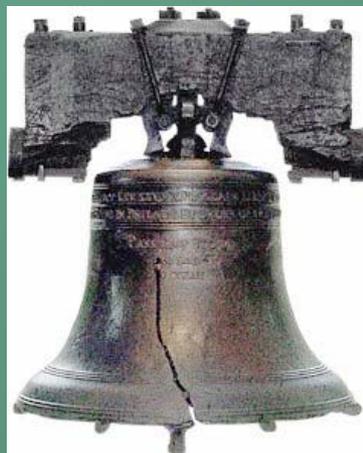
INTENSITY	I	II-III	IV	V	VI	VII	VIII	IX	X+
SHAKING	Not felt	Weak	Light	Moderate	Strong	Very strong	Severe	Violent	Extreme
DAMAGE	none	none	none	Very light	Light	Moderate	Moderate/Heavy	Heavy	Very Heavy

Modified Mercalli Intensity

- I Barely felt
- II Felt by only few people
- III Felt noticeably, standing autos rock slightly
- IV Felt by many, windows and walls creak
- V Felt by nearly everyone, some dished and windows broken
- VI Felt by all, damaged plaster and chimneys
- VII Damage to poorly constructed buildings
- VIII Collapse of poorly constructed buildings,
slight damage to well built structures
- IX Considerable damage to well constructed buildings,
buildings shifted off foundations
- X Damage to well built wooden structures, some masonry
buildings destroyed, train rails bent, landslides
- XI Few masonry structure remain standing, bridges
destroyed, ground fissures
- XII Damage total

Normal Modes

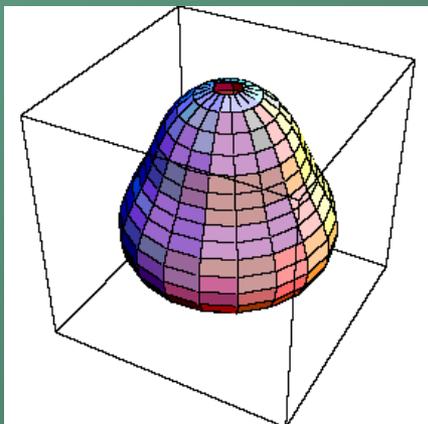
Liberty Bell
(USA)



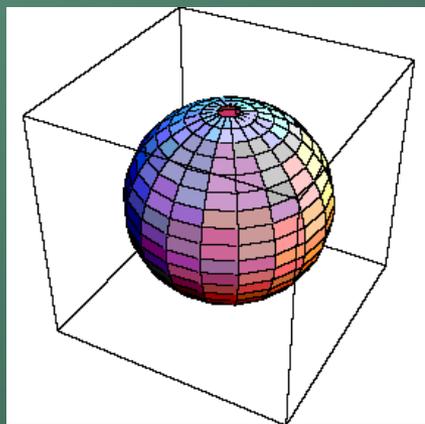
Useful for studies of

- Interior of the Earth
- Largest earthquakes

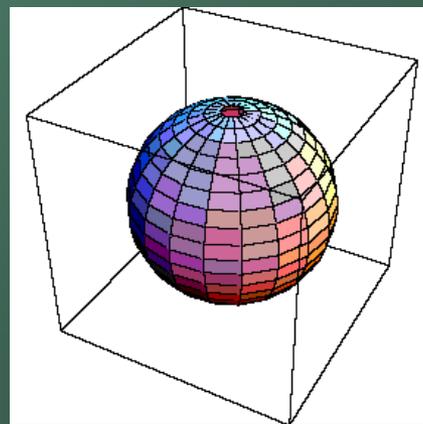
$l=1$ $m=1$



$l=1$ $m=2$

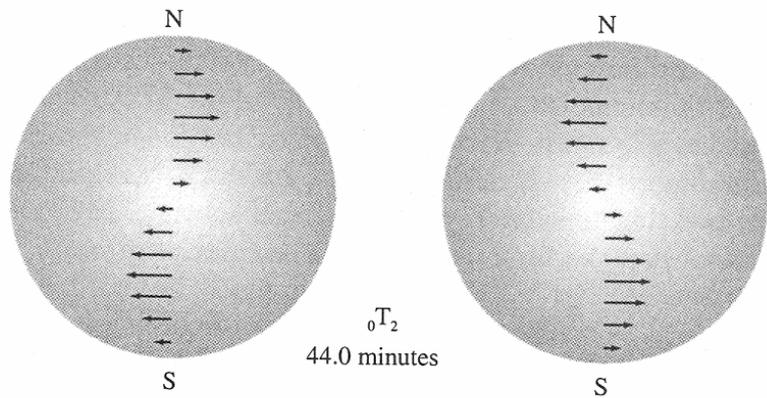


$l=1$ $m=3$

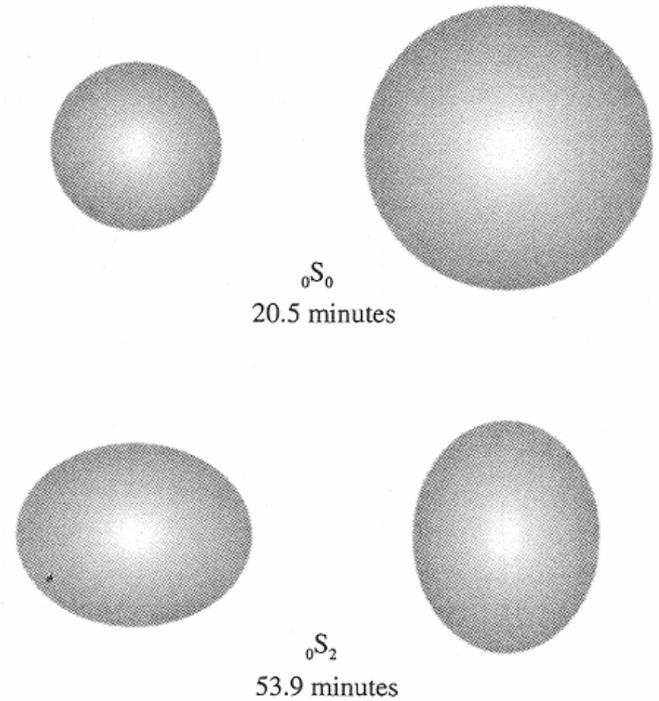


Houseman <http://earth.leeds.ac.uk/~greg/?Sphar/index.html>

Toroidal and Spheroidal Modes



Toroidal



Spheroidal

Natural Vibrations of the Earth

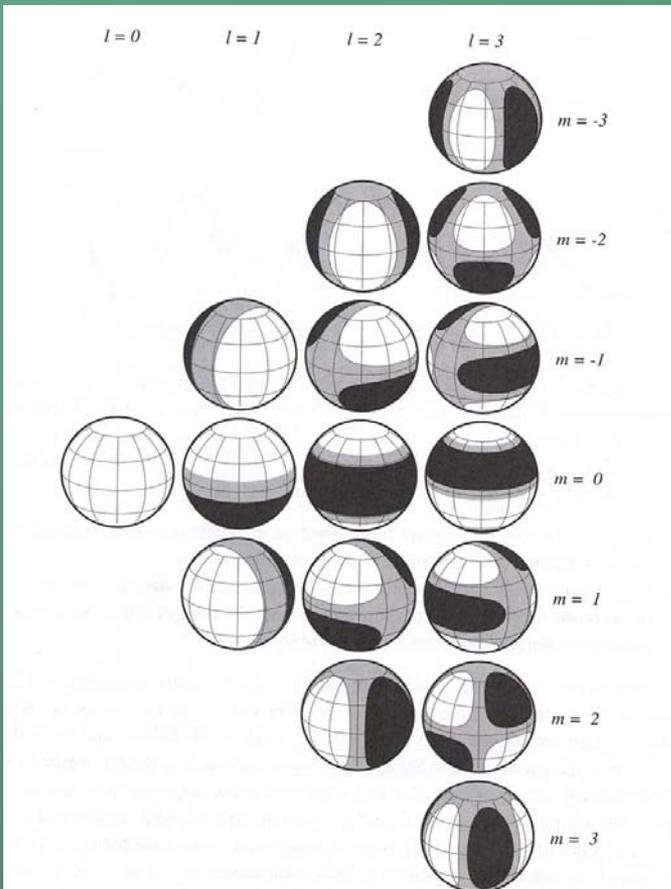


TABLE 4.2 Some Observed Normal-Mode Periods

Spheroidal modes	T (s)	Toroidal modes	T (s)
${}_0S_0$	1227.52	${}_0T_2$	2636.38
${}_0S_2$	3233.25	${}_0T_{10}$	618.97
${}_0S_{15}$	426.15	${}_0T_{20}$	360.03
${}_0S_{30}$	262.09	${}_0T_{30}$	257.76
${}_0S_{45}$	193.91	${}_0T_{40}$	200.95
${}_0S_{60}$	153.24	${}_0T_{50}$	164.70
${}_0S_{150}$	66.90	${}_0T_{60}$	139.46
${}_1S_2$	1470.85	${}_1T_2$	756.57
${}_1S_{10}$	465.46	${}_1T_{10}$	381.65
${}_2S_{10}$	415.92	${}_2T_{40}$	123.56

Summary

Rays

Earth structure causes complicated ray paths through the Earth (P, PKP, PcP)

Wave theory explains

- P and S waves
- Static displacements
- Surface waves

Normal Modes

The Earth rings like a bell at long periods